

## COMMONWEALTH OF PENNSYLVANIA

Robert P. Casey, Governor

## DEPARTMENT OF ENVIRONMENTAL RESOURCES Arthur A. Davis, Secretary

OFFICE OF PARKS AND FORESTRY Patrick J. Solano, Deputy Secretary

### BUREAU OF TOPOGRAPHIC AND GEOLOGIC SURVEY Donald M. Hoskins, Director



# CONTENTS



The Fourth Pennsylvania Geological Survey at 75: 1919 to 1994	1
Hooked on the rocks: exploring the geology of the Youghiogheny River	
Lake—a boating field trip	2
Sedimentary evaporites in the Gatesburg Formation of central Pennsyl-	
vania 1	10
Earth Science Teachers' Corner 1	16

# **ON THE COVER**

The "stair-step" exposure of the Lower Conemaugh Pine Creek marine zone along the shoreline of the Youghiogheny River Lake, Confluence, Pa., locality 11 (see article on page 2). The outcrop was exposed due to the lowering of the lake level. Photograph by W. E. Kochanov.

# PENNSYLVANIA GEOLOGY

PENNSYLVANIA GEOLOGY is published quarterly by the Bureau of Topographic and Geologic Survey, Pennsylvania Department of Environmental Resources, P. O. Box 8453, Harrisburg, PA 17105–8453.

Editors: Christine M. Dodge, Anne B. Lutz, and Donald M. Hoskins.

Contributed articles are welcome; for further information and guidelines for manuscript preparation, contact D. M. Hoskins at the address listed above.

Articles may be reprinted from this magazine if credit is given to the Bureau of Topographic and Geologic Survey.

VOL. 25, NO. 2

**SUMMER 1994** 

# The Fourth Pennsylvania Geological Survey at 75: 1919 to 1994

The present (Fourth) Topographic and Geologic Survey of Pennsylvania was created on June 7, 1919, when then Governor Sproul signed Act 205 into law. Although Governor Sproul's appointment letter to the first employee of the new Bureau was dated September 4, 1919, the work of the survey actually began on September 1, reflecting the urgent need of the day. On that date, Dr. George H. Ashley, widely respected for a lengthy career in both state and federal geological surveys, became State Geologist. He immediately began to deal with an accumulation of problems in Pennsylvania that required both an understanding of geology and the support of accurate topographic maps. His successors as State Geologists have been Ralph Stone, Stanley Cathcart, Carlyle Gray, Arthur Socolow, and myself.

During the succeeding 75 years, the staff of the Topographic and Geologic Survey of Pennsylvania has served the Commonwealth by providing answers to problems that require an understanding of the geology and topography of our state. The Bureau has served Pennsylvania citizens by responding to uncounted thousands of requests for specific geological and topographic information and for explanations of the many geological features and phenomena present in our Commonwealth. In order to provide reliable scientific information supporting these responses, the Bureau has independently conducted or assisted in extensive field investigations in all areas of the Commonwealth. These investigations include the collection of geologic data concerning the location and nature of Pennsylvania's rocks and surficial materials and their energy, mineral, and water resources. The data have been analyzed, synthesized, and provided to the public through detailed, as well as popularly written, reports and maps. The number of reports and maps issued since 1919 now totals nearly 900.

These accomplishments are the result of 75 years of work by dedicated and competent geologists and support staff of the Bureau, for whom public service has a high personal value, following the principles established by Dr. Ashley. Their work and the products of the Bureau have been complemented by the efforts of the larger community of geologists who investigate and analyze the complex geology of

(continued on page 15)

# HOOKED ON THE ROCKS: Exploring the Geology of the Youghiogheny River Lake—A Boating Field Trip

## by James R. Shaulis Pennsylvania Geological Survey

You are out fishing one morning on the Youghiogheny River Lake. As the sun begins to rise, you make your first cast of the day. Trying to drop your lure near the steep rocky shore, you accidentally overthrow. Your lure becomes wedged in a rock crevasse so tightly that you are forced to go ashore to free it. When you bend down to retrieve it, you notice that some of the rock layers contain many different types of marine fossils. Other rock slabs that have slid down from above have rippled surfaces, and a few are also marked with sinuous wormlike trails. When you look up to see where these rocks came from, you notice that higher up on the shoreline, the rocks change in color from pale olive to gravish red. You have many questions about what you see here. If the answers could be found on a boating excursion that included beautiful scenery, fishing, fossil hunting, and a basic introduction to the various types of sedimentary rocks and the geologic history of western Pennsylvania, would that be a catch worth keeping? If so, cast your eves on this article and reel it in.

THE CREATION, CARE, AND MAINTENANCE OF AN OUTCROP.

The Youghiogheny River Lake (YRL) (Figure 1) was completed by the U.S. Army Corps of Engineers in 1943 in conjunction with the Flood Control Act of 1938. The YRL provides flood protection and also is a water supply for downstream communities and commerce. Its water level is adjusted continuously by the Corps of Engineers according to a schedule that takes into account seasonal precipitation amounts in the YRL watershed. On the average, the level is adjusted by 0.5 foot per day up or down over a 50-foot range between 1,390 feet and 1,440 feet above mean sea level. However, during very wet or very dry periods, the range can increase by 20 feet in either direction. The right-hand column of Table 1 contains a listing of water levels by month for an average precipitation profile based on records for the past 30 years.

Before the YRL existed, the Youghiogheny River cut a deep gorge-like valley hundreds of feet deep across the landscape of this area. Field notes of geologists who mapped in the region indicated that few natural rock exposures existed, except for some resistant sandstone ledges. Most of the rocks in this area were covered by thick vegetation, soil cover, and talus.

However, since the construction of the lake, rocks have become exposed, especially on the steeper outside cut banks of the former Youghiogheny River. The continuous raising and lowering of the water level has effectively stripped

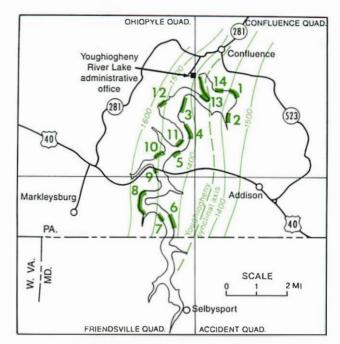


Figure 1. Map showing the location of rock exposures along the Youghiogheny River Lake shoreline and structure contours on the basal coquina limestone of the Pine Creek marine zone.

away the soil and vegetation from those areas to expose the rocks. Once exposed, the rocks are swept clean of sediment by wind and powerboat-generated wave action. The natural slope and the daily incremental change in water level has created a stairstep outcrop pattern, which provides many opportunities for examining beds in a three-dimensional perspective (see cover photograph). During the spring and summer months, the rocks are completely submerged; in late summer to late fall, they are exposed and washed clean before being subjected to many freeze/thaw cycles during the beginning of winter. This yearly sequence of events has greatly accelerated the weathering of these rocks and has kept the outcrops uncovered and fresh like a rocky ocean shoreline, so that today, with the combined efforts of man and nature, miles of pristine rock exposures exist where none had been before. Figure 1 shows the locations of the 14 best-developed outcrops on the YRL shoreline in Pennsylvania.

**OBSERVATIONS AND INTERPRETATIONS OF THE ROCKS.** The YRL lies completely within the Youghiogheny synclinal structure. The axis of this syncline trends northeast-southwest and intersects

						LU	cality	numbe	51						Lake
No.	1	2	3	4	5	6	7	8	9	10	11	12	13	14	elev."
J	n pc	n pc	n pc	n pc	рс	n pc	b	m	b	b	n pc	b	wr	wr	1,400
F	n pc	n pc	n (pc)	n pc	pc	n pc	b	m	b	b	n pc	b	wr	wr	1,410
M	n (pc)	n	n	n pc	pc²	n (pc)	b	m	b	b	n pc	b	wr	wr	1,420
A		n²	Π²		pc <sup>2</sup>	n²	b <sup>2</sup>	m²	b		n <sup>2</sup>				1,437
M		U <sub>5</sub>	Π²		pc <sup>2</sup>	n²	b <sup>2</sup>	m²	b		IL5				1,440
J			n²		pc <sup>2</sup>	n²	b <sup>2</sup>	m²	b		n²				1,440
J	п	п	n		pc	n²	b <sup>2</sup>	m²	b		<b>П</b> <sup>2</sup>				1,438
Α	п	n	n		pc	п	b	m	b	(b)	n				1,433
S	(pc) n	n	(pc) n	(pc) n	рс	(pc) n	b	m	b	b	n pc	b	wr	wr	1,421
0	n pc	n pc	n pc	n pc	pc	n pc	b	m	b	b	n pc	b	wr	wr	1,407
N	n pc	n pc	n pc	n pc	pc	n pc	b	m	b	b	n pc	b	wr	wr	1,394
D	n pc	n pc	n pc	n pc	pc	n pc	b	m	b	Ь	n pc	Ь	wr	wr	1,395

the shoreline of the YRL at locality 13 (Figure 1). Its southeastern limb can be seen dipping gently 1 degree to the northwest toward the axis at locality 1. The dip on the northwestern limb varies from 1 to 2.5 degrees to the southeast but is hard to see because most of the outcrops on this side of the syncline are oriented parallel to strike.

The exposed stratigraphic sequence totals about 230 feet along the YRL shoreline and contains four discrete intervals: (1) the Mahoning coal and sandstone interval; (2) the Brush Creek coal and marine zone interval; (3) the Pine Creek and Nadine marine zone interval; and (4) the Lower Bakerstown coal and Woods Run marine zone interval (Figure 2). Composite columnar sections and interpretations of depositional environments for each interval are shown in Figure 3. The depositional history for each interval is briefly discussed below, in ascending stratigraphic order.

A. Mahoning Coal and Sandstone Interval (Locality 8). The Mahoning coal and sandstone interval is the only nonmarine sequence occurring along the lake. It makes up the longest continuous outcrop, nearly 1 mile in length.

After a river that was flowing across a deltaic plain deposited the Lower Mahoning sandstone, it migrated out of the area. A slow

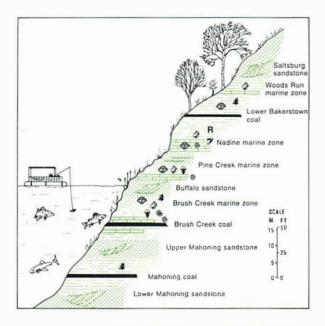


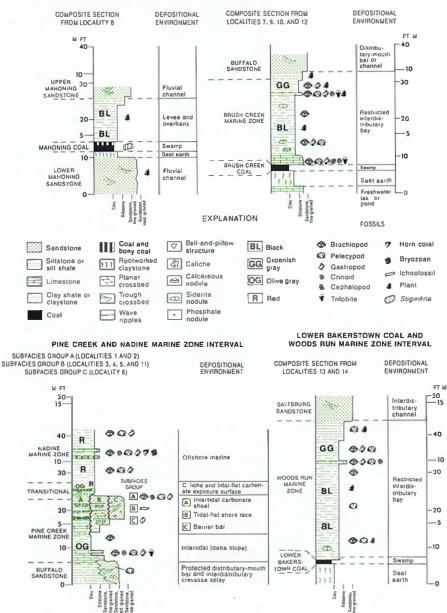
Figure 2. Composite columnar section of Upper Pennsylvanian rocks (Glenshaw Formation, Lower Conemaugh Group) exposed along the YRL shoreline. The sketch shows rock units exposed at locality 1 in mid-September in an average rainfall year. The fish under the boat are rock bass and freshwater sandstone sharks. See Figure 3 for explanation of symbols and patterns.

subsidence of the land then occurred, and large, shallow, poorly drained depressions developed. This provided an ideal setting for plant growth and the development of a peat swamp that eventually became the Mahoning coal. The swamp was then covered with clay, organic debris, and silt from overbank deposits of a nearby fluvial system, Finally, another river meandered back into the area and deposited the Upper Mahoning sandstone at the top

of the section. A generalized diagram illustrating depositional environments common to a Pennsylvanian delta is shown in Figure 4. Of special interest is the upper portion of the Mahoning coal seam, which contains well-preserved sideritized *Stigmaria* fossils up to 0.8 inch wide and 3.0 feet long.

**B.** Brush Creek Coal and Marine Zone (Localities 7, 9, 10, and 12). The Brush Creek coal and marine zone is exposed at four separate localities on the western shoreline of the YRL.

The freshwater limestone at the base of the section is interpreted to have been deposited in a freshwater pond or lake. As this lake filled with lime mud and fine noncarbonate sediments, the water shallowed, allowing plant growth, and the area evolved into a swamp that accumulated peat, which eventually formed the Brush Creek coal. A rise in sea level led to a marine transgression over the swamp, transforming the area into a restricted interdistributary bay. In this environment, dark fossiliferous shales and thin coquina limestones containing a nearshore marine fauna were deposited. Land was close by, as indicated by the presence of crinoid columnals and



#### MAHONING COAL AND SANDSTONE INTERVAL

BRUSH CREEK COAL AND MARINE ZONE INTERVAL

Figure 3. Composite columnar sections of rock exposures found along the Youghiogheny River Lake and their interpreted depositional environments.

intact fern fronds together in one of the coquina limestone beds. As the interdistributary bay continued to fill with sediment and prograde seaward, a coarsening-upward sequence of rocks was deposited that culminated with the Buffalo sandstone.

C. The Pine Creek and Nadine Marine Zones (Localities 1, 2, 3, 4, 5, 6, and 11). The Pine Creek-Nadine marine zone interval is exposed at seven different localities covering more than 2 miles of the YRL shoreline.

The sediments underlying the Pine Creek were formed either in an interdistributary bay environment by a distributary-mouth bar (locality 1

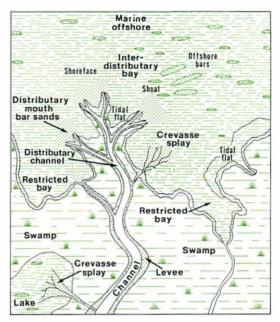
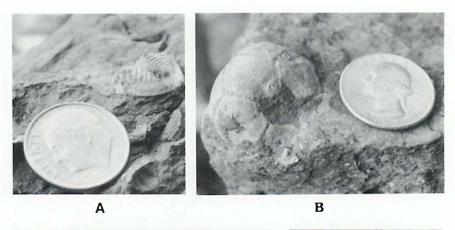


Figure 4. Generalized diagram showing depositional environments common to a Pennsylvanian delta (modified from Flores and Arndt, 1979). See Figure 3 for explanation of patterns.

only) or as crevasse-splay deposits. A subsequent rise in sea level resulted in a marine transgression. The coquina limestone and fossiliferous shales (Figures 5A, 5C, and 5D) at the base of the Pine Creek represent the initial transgressive phase. Following this phase, a shallowing of the water occurred as the land prograded seaward. This is reflected in a gradual coarsening-upward pattern in the Pine Creek. However, in the upper portion, this trend is also accompanied by complex lateral facies changes in what is interpreted as a nearshore intertidal environment. Within this intertidal setting, the following subfacies groups have been identified: intertidal carbonate shoal (group A), tidal-flat shore face (group B), and barrier bar (group C)(Figure 3).

Following the deposition of the Pine Creek in a nearshore intertidal environment, a drying out occurred, as indicated by the appearance of caliche beds. After subaerial exposure of the land, marine conditions returned and resulted in the deposition of a red clay shale interbedded with thin limestones, called the Nadine marine zone. The presence of crinoids (Figure 5B), corals, bryozoans, and phosphate nodules in the limestones and brachiopods in the red shales, and



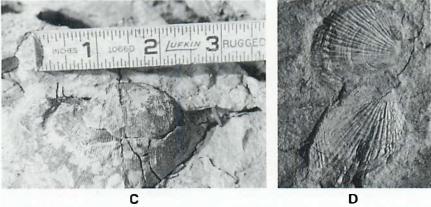


Figure 5. Photographs of fossils found in the rocks along the YRL shoreline. A, Trilobite pygidium in basal coquina limestone of the Pine Creek marine zone from locality 5. B, Crinoid calyx found in the Nadine limestone at locality 3. C, Large "adult" *Juresania* sp., a brachiopod from the basal Pine Creek at locality 11. D, *Dunbarella* sp., a brachiopod from the lower Pine Creek shale at locality 11. The length of the large valve is 25 mm.

the lack of any coarse-grained siliciclastics in the Nadine interval indicate that it was deposited in an offshore marine setting. The red color of the clay shale may have resulted from the erosion of a reddish, iron-rich soil into the marine depositional environment.

**D.** The Lower Bakerstown Coal and Woods Run Marine Zone (Localities 13 and 14). This interval is lithologically nearly identical to the Brush Creek coal and marine zone and is interpreted to have a similar depositional history (Figure 3).

**PLANNING YOUR VISIT TO SEE THE SITES.** Before planning to visit any of the outcrop localities along the lake, first check the exposure chart shown in Table 1 to see which sites are exposed at that time of year. Better yet, contact the Army Corps of Engineers at YRL to find out if the current lake level corresponds to the level shown on the chart. Otherwise, determine what sites are accessible by finding the month on the chart in which the lake level is approximately the same as the current lake level. The best time to examine all of the outcrops is in late summer to late fall when water levels are lowest.

To get to the outcrops, pontoon boats are available for daily rental at the Youghiogheny Lake Marina from June 1 through the middle of September. However, all outcrops are accessible by car and foot for the truly determined. Before attempting to do this, consult the 7.5minute-quadrangle map of the area. Many roads that encircle the lake are narrow semiprivate routes having limited off-road parking. If you are near a private residence, it is advisable to ask permission before leaving your car parked for an extended period. WARNING: The lake is surrounded by steep slopes and cliffs that are made up of highly weathered, incompetent rocks. Vertical drops of more than 20 feet occur along the shoreline in many areas. The author can assure you that access via this route can be very quick, but very painful, so use extreme caution if you go down to the shoreline on foot in these areas.

THE END OF THE LINE. The sun has now almost set, but it seems as though only an hour has passed. You have visited ancient oceans, coastlines, bays, rivers, and swamps without leaving the shoreline of the Youghiogheny River Lake. You have marveled at how the natural geologic processes and man's presence have changed and continue to change the landscape over time.

When you finally arrive back home, a family member comments on how the fishing must have been good today because you are home so late. Not having any fish to show off, you reply, "Well you see, what happened is that, while I was fishing, I got hooked on some rocks and I just couldn't get loose."

## REFERENCE

Flores, R. M., and Arndt, H. H. (1979), Depositional environments of Middle Pennsylvanian Series in proposed Pennsylvanian System stratotype, in Englund, K. H., and others, eds., Proposed Pennsylvanian System stratotype, Virginia and West Virginia, Ninth International Congress of Carboniferous Stratigraphy and Geology, American Geological Institute Selected Guidebook Series 1, p. 115–121.

# Sedimentary Evaporites in the Gatesburg Formation of Central Pennsylvania

by Christopher D. Laughrey Pennsylvania Geological Survey

> "It's these changes in latitude... nothing remains quite the same." —Jimmy Buffet, 1976

**INTRODUCTION.** The winter of 1993–94, with no fewer than 16 snowstorms, caused many a Pennsylvanian to dream about balmy southern climes and tropical shores. While shovelling through the drifts left in my driveway by a hydrologic cycle gone berserk, I tried to find consolation in the fact that the Keystone State once lay near the equator. A little more than 500 million years ago, Pennsylvania resembled the Bahamas, although it was probably much drier. Drifts of sand rather than snow lined the shore, and often, much of the water available there evaporated, leaving behind only a mineral legacy.

THE GATESBURG FORMATION AND "GREAT AMERCAN BANK." Rocks of Late Cambrian age in the Gatesburg Formation of central Pennsylvania are part of a thick regional succession of platform carbonates that accumulated on the southwest-facing coast of the ancient continent called Laurentia, which was positioned athwart the equator (see back cover). The Cambrian and Ordovician Periods were times of widespread carbonate deposition in what would later become North America. Broad carbonate shelves extended along the lengths of what would become the Appalachians in the east and the Rockies in the west (Hardie, 1986, p. 40). In the area of the future Appalachians, this Cambrian and Ordovician platform produced a wedge of carbonate rocks more than 1,860 miles long, more than 186 miles wide, and as much as 2.2 miles thick (Colton, 1970)! It is one of the largest accumulations of carbonate sediments in the geologic record (Hardie, 1986), an accumulation that the eminent carbonate geologist R. N. Ginsburg (1982) christened the "Great American Bank" (similar to the modern Great Bahama Bank).

In central Pennsylvania, the Gatesburg Formation consists of repeated (cyclic) sequences of peritidal carbonates, noncyclic subti-

dal carbonates, and two rather prominent quartz-sandstone zones. The cyclic peritidal carbonates exhibit well-defined shallowing-upward tidal-flat sequences, as schematically illustrated in Figure 1.

**ANCIENT CLIMATES.** Paleogeographic and paleo-oceanographic placement of the Laurentian landmass indicates that the climate was probably semiarid during deposition of the Upper Cambrian rocks of central Pennsylvania. Because the southwest-facing coast of Laurentia lay between 0 and 30 degrees south latitude (see back cover), the dominant wind pattern likely was a southeasterly trade wind that carried warm maritime air from the moist western side of a high-pressure center in the lapetus Ocean (Riley and others, 1993, p. 11–12). This air movement would have provided abundant rainfall along the east coast of Laurentia, but the southwest coastal climate would have been relatively dry and similar to those of modern-day Morocco, Baja California, and western Australia (Wilde, 1991; Riley and others, 1993).

The geological evidence for a semiarid setting is mostly indirect. Shallowing-upward tidal-flat sequences in the Gatesburg Formation contain several indicators of subaerial exposure and desiccation, such as mudcracks and intraclast chips in association with isolated, displacive nodules of sparry calcite. Hardie (1986) and Read (1989) interpreted the calcite nodules as former anhydrite nodules formed in the supratidal zone.

**EVAPORITES.** The Pennsylvania Geological Survey recently participated in a two-year multidisciplinary research program designed to measure and predict petroleum reservoir heterogeneity in Upper Cambrian rocks in Pennsylvania and Ohio (Riley and others, 1993). During this work, Survey geologist John Harper and I were surprised and delighted to find sedimentary evaporites in core samples of the Gatesburg Formation.

The Shell Oil Company drilled the Shade Mountain #1 well in Fayette Township, Juniata County, in 1964. One of the cores recovered from this well consists of almost 26 feet of dolostone and a 4-foot-thick sequence of sedimentary anhydrite near the top (Figures 1B and 2). The anhydrite occurs between depths of 10,013 and 10,017 feet. The lowest part of the anhydrite zone in the core is transitional with subjacent oolitic, fossiliferous, mottled, and burrowed dolostone.

The bulk of the evaporite in the core consists of moderateyellowish-brown, slightly dolomitic anhydrite. The anhydrite displays wavy, anastomosing, and laminated bedding. It also contains very thin, discontinuous and wispy, nonparallel laminations of finely crystalline dolomite intercalated with micron-sized, subhedral crystals of

	MUDCRACKED FLAT LAMINITE	SUPRATIDAL
	PRISM-CRACKED WAVY LAMINITE	UPPER = INTERTIDAL ALGAL FLATS
	RIBBON ROCK	LOWER INTERTIDAL MIXED SAND- MUD FLATS
	CROSS – STRATIFIED GRAIN STONES	SUBTIDAL = OOID-PELOID SAND SHOALS
	THROMBOLITIC BLOHERMS	ALGAL = PATCH REEFS
The analyze	FLAT PEBBLE CONGLOMERATE	= STORM SHEET DEPOSITION

Α

9

STRATIGRAPHIC UP

CORE DEPTH

Figure 1. Shallowing-upward carbonate sequences are stratigraphic successions in which each unit was deposited in progressively shallower water. These successions are repeated many times within a stratigraphic section of shallow-water carbonate rocks. A, Lithologies and interpreted depositional environments of several sedimentary subfacies in shallowingupward sequences found in Cambrian and Ordovician rocks in the central Appalachians (from Hardie, 1986). B, Graphical core description of the Gatesburg Formation in the Shade Mountain well, Juniata County (modified from Riley and others, 1993). The core contains repeated shallowingupward sequences. Note the presence of anhydrite near the top of the core.

#### LITHOLOGIES

	10.010	SUBTIDAL SHELF-LAGOON	DOLOSTONE, mottled and burrowed			
~	<u>Å^^^ ^ ^ </u>		NODULAR ANHYDRITE containing dolomite			
	10,015	banks having sabkha deposits	LAMINATED EVAPORITES contain- ing dolomile and stromatolites			
		bank lagoon or "pond" deposits	DOLOSTONE. oolitic, fossililerous, mottled, and burrowed			
		SUBTIDAL TO INTERTIDAL (?) shoal or levee bank and shoal or channel flank	OOID DOLOGRAINSTONE, grades down to MIXED OOID-PELOID DOLOGRAINSTONE and DOLO- PACKSTONE, and PELOID-MIXED FOSSIL DOLOPACKSTONE			
		SUBTIDAL SHELF-LAGOON	MIXED FOSSIL DOLOPACK- STONES AND DOLOWACKE- STONES			
		SUBTIDAL SHOAL OR CHANNEL FLANK	OOID-PELOID DOLOGRAINSTONE, quartzose, fossiliferous; contains rip-up clast conglomerates			
	10.035	SUBTIDAL SHELF-LAGOON	DOLOSTONE, mottled and bioturbated			
		EXPLANATION				
	Doto	stone .I. Dolomilic	Peloids			
	Calc	areous dolostone 🛛 📉 Nodular beddi	ing Stromatolites			

Laminated bedding

Ooids
B

DEPOSITIONAL SETTING

 $\Lambda \Lambda \Lambda$ 

Anhydrite

Quartz sand

anhydrite (Figure 3). Coarser anhydrite in this part of the sequence consists of subhedral crystals ranging in size from 0.03 to 0.75 mm in diameter (Figure 3).

The upper part of the anhydrite sequence in the Shade Mountain well core consists of light-olive-gray to yellowish-gray nodular anhydrite. The change from laminated to nodular bedding is transitional, and a zone of decreasing lamination and crystal size occurs between the laminated and nodular intervals. The laminae in the transitional zone are indistinct, and some nodules appear within the laminae. Distinct nodules in the upper part of the sequence consist of aggregates of subhedral anhydrite crystals in a groundmass of micron-sized dolomite. The latter contains the microfossils *Epiphyton* (a calcareous red alga) and *Renalcis* (a cyanophyte, or "blue-green

alga"). These remnants of mound-building organisms are common in Cambrian and Ordovician rocks of the central Appalachians.

The anhydrite in the Shade Mountain well core exhibits two principal habits, laminated and nodular. Nodular anhydrite is a common habit in both modern and ancient evaporite deposits, and its presence is a key factor in the recognition of paleosabkhas in the rock record (Tucker, 1981; Kendall, 1984). Laminated anhydrite has been variously interpreted as basinal varves, algal mats, or coalescing nodules. Nodular and laminated anhydrite commonly occur together, and both form subaqueously in shallow or deep standing water (Dean and others, 1975). The sequence in the Shade Mountain well core occurs within well-defined cyclic peritidal deposits and

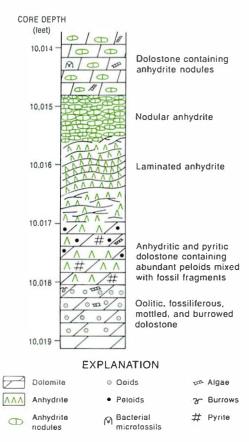


Figure 2. Detailed description of the anhydrite interval from near the top of the core shown in Figure 1.

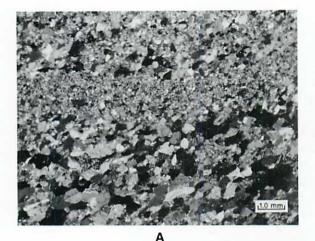


Figure 3. Thin-section photomicrographs of the laminated anhydrite in the Shade Mountain well core. A, Low magnification reveals the laminated habit of the anhydrite. B, Higher magnification shows the crystal texture in the coarser laminae.

can be confidently interpreted as having been produced on a tidal flat (Figure 1). The evaporites in the Shade Mountain core probably formed along the high portion of a carbonate bank or in interbank ponds and lagoons as subaqueous evaporites.

Although the specific mechanism of



В

original evaporite precipitation and subsequent diagenetic alteration of these beautiful anhydrites in the Shade Mountain core remain to be proven, the fact that unambiguous sedimentary evaporite minerals do occur in Upper Cambrian rocks of central Pennsylvania helps to support published paleoclimatic interpretations of ancient continents. This occurrence of anhydrite in the Gatesburg Formation provides direct evidence of at least semiarid, and perhaps actual arid, conditions during Late Cambrian time on the southwest-facing coast of Laurentia.

### REFERENCES

Bambach, R. K., Scotese, C. R., and Ziegler, A. M. (1980), *Before Pangea: the geog*raphies of the Paleozoic world, American Scientist, v. 68, p. 26–38. Colton, G. W. (1970), *The Appalachian basin—its depositional sequences and their geologic relationships*, in Fisher, G. W., and others, eds., *Studies of Appalachian geology: central and southern*, New York, Interscience Publishers, p. 5–47.

Dean, W. E., Davies, G. R., and Anderson, R. Y. (1975), Sedimentological significance of nodular and laminated anhydrite, Geology, v. 3, p. 367–372.

Ginsburg, R. N. (1982), Actualistic depositional models for the Great American Bank (Cambro-Ordovician) [abs.], International Congress on Sedimentology, 11th, Hamilton, Ontario, Canada, Abstracts, International Association of Sedimentologists, p. 114.

Hardie, L. A. (1986), Ancient carbonate tidal-flat deposits, in Hardie, L. A., and Shinn, E. A., Carbonate depositional environments, modern and ancient—Part 3: tidal flats, Colorado School of Mines Quarterly, v. 81, no. 1, p. 37–57.

Kendall, A. C. (1984), *Evaporites*, in Walker, R. G, ed., *Facies models*, 2nd ed., Geoscience Canada, Reprint Series 1, p. 259–296.

Read, J. F. (1989), Controls on evolution of Cambrian-Ordovician passive margin, U.S. Appalachians, in Crevello, P.D., and others, eds., Controls on carbonate platform and basin development, Society of Economic Paleontologists and Mineralogists Special Publication 44, p. 147–165.

Riley, R. A., and others (1993), Measuring and predicting reservoir heterogeneity in complex deposystems: the Late Cambrian Rose Run Sandstone of eastern Ohio and western Pennsylvania, Appalachian Oil and Natural Gas Research Consortium, National Research Center for Coal and Energy, West Virginia University, Morgantown, 257 p.

Tucker, M. E. (1981), *Sedimentary petrology, an introduction*, New York, John Wiley and Sons, 252 p.

- Warren, J. K., and Kendall, C.G. (1985), Comparison of sequences formed in marine sabkha (subaerial) and salina (subaqueous) settings—modern and ancient, American Association of Petroleum Geologists Bulletin, v. 69, p. 1013–1023.
- Wilde, Pat (1991), Oceanography in the Ordovician, in Barnes, C. R., and Williams, S. H., eds., Advances in Ordovician geology, Geological Survey of Canada Paper 90–9, p. 283–298.

The Fourth Pennsylvania Geological Survey at 75: 1919 to 1994 (continued from page 1)

Pennsylvania. Although the tasks ahead will be even more demanding as increased use is made of Pennsylvania's land surface and earth resources, on this occasion of the Survey's 75th anniversary, I commend the past and present geologists, hydrogeologists, and technical and administrative assistants of the Pennsylvania Bureau of Topographic and Geologic Survey for jobs very well done!

Donald M. Hollins

Donald M. Hoskins State Geologist

# EARTH SCIENCE TEACHERS' CORNER

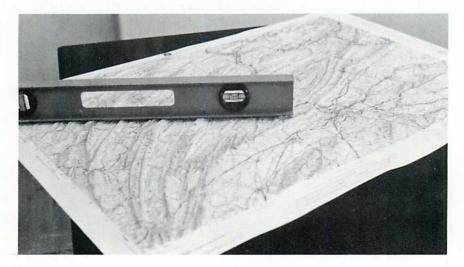
# The Carpenter's Level as a Geomorphological Teaching Tool

## by William Kreiger

Physical Sciences Department, York College York, PA 17405–7199

As part of an introductory laboratory mapping exercise that utilizes road maps, topographic maps, and raised-relief maps of Pennsvlvania, the students observe the mountains in the Ridge and Valley physiographic province on all maps. After they look at the ridges on the "flat" maps, they place a carpenter's level on the mountaintops in both the northsouth and northwest-southeast directions at several locations on the Hubbard NK18-10 Harrisburg raised-relief map, which has a scale of 1:250,000. The students

find that the level indicates that the tops of the mountains are all at about the same elevation. The bubble indicates that these tops form a level surface. Even though the students can see from the topographic maps that the elevations are quite close, the simple act of placing a carpenter's level on the raised-relief map does wonders for their comprehension. This new awareness opens the door for discussion of peneplains versus rock resistance and other aspects of erosion that produced this topography.



## PENNSYLVANIA GEOLOGICAL SURVEY STAFF

Donald M. Hoskins, Bureau Director

## DIRECTOR'S OFFICE

717-787-2169

Access no. for the hearing impaired: 1-800-654-5984 (TDD) Pa. AT&T relay service

#### Administrative Services

Shirley J. Barner, *Clerical Supervisor* Sharon E. Garner, *Clerk Typist* Jody R. Smith, *Clerk Typist*  Administrative Services (Cont.) Janet L. Wotring, Administrative Assistant

### GEOLOGIC AND GEOGRAPHIC INFORMATION SERVICES DIVISION

717-787-8162 Christine M. Dodge, Division Chief

Geologic Library Richard C. Keen, Librarian

Geologic Information Section Anne B. Lutz, *Geologist* Caron E. O'Neil, *Geologist* Thomas G. Whitfield, *Geologist* 

#### **Geographic Information Section**

James H. Kaiser, *Engineering Technician* William H. Stoner, *Engineering Technician* 

#### **Geologic Drafting Section**

John G. Kuchinski, Section Chiel Lajos J. Balogh, Cartographic Drafter (Pittsburgh Office) James H. Dolimpio, Cartographic Drafter John G. Petro, Cartographic Drafter (Pittsburgh Office)

### **GEOLOGIC MAPPING DIVISION**

717–787–6029 Jon D. Inners, Division Chief

### Eastern Regional Section

Rodger T. Faill, Section Chief David B. MacLachlan, Geologist William D. Sevon, Geologist

Geotechnical Section Helen L. Delano, *Geologist* William E. Kochanov, *Geologist* 

#### Western Regional Section

Albert D. Glover, Section Chief Clifford H. Dodge, Geologist Leonard J. Lentz, Geologist John C. Neubaum, Geologist James R. Shaulis, Geologist Viktoras W. Skema, Geologist

### **GEOLOGIC RESOURCES DIVISION**

717–787–5828 Samuel W. Berkheiser, Division Chief

### Groundwater Geology Section

Michael E. Moore, *Section Chief* Thomas A. McElroy, *Hydrogeologist* Dawna S. Yannacci, *Hydrogeologist* 

Mineral Resources Section Robert C. Smith, II, Section Chiel John H. Barnes, Geologist Leslie T. Chubb, Laboratory Technician

Water Well Drillers Licensing and Records Section Donna M. Snyder, Section Chief Mari G. Barnhart, Clerk

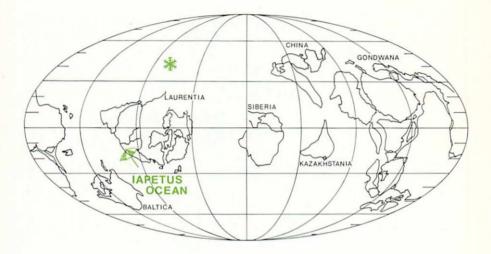
#### Subsurface Geology Section

500 Waterfront Drive Pittsburgh, PA 15222–4745 412–442–4235 John A. Harper, Section Chiel Cheryl L. Cozart, Descriptive Statistician Kathy J. Flaherty, Geologist Joseph E. Kunz, Clerk Typist Christopher D. Laughrey, Geologist Lynn J. Levino, Clerk Typist Antonette K. Markowski, Geologist (Harrlsburg Office) Joseph R. Tedeski, Geologist

IN COOPERATION WITH THE U.S. GEOLOGICAL SURVEY TOPOGRAPHIC MAPPING GROUNDWATER-RESOURCE MAPPING **GLOBAL GEOGRAPHY IN THE LATE CAMBRIAN** 

(modified from Bambach and others, 1980)

(see article on page 10)



Approximate position of Pennsylvania on the southwest coast of Laurentia in the Late Cambrian

Rosition of Pennsylvania today

Bureau of Topographic and Geologic Survey Department of Environmental Resources P. O. Box 8453 Harrisburg, PA 17105–8453

Buik Rate U.S. Postage PAID Harrisburg, PA Permit No. 747

Address Corrections Requested

